

مرکز تحقیقات آبهای زیرزمینی (متآب)
Groundwater Research Center (GRC)

هیدروژئوشیمی و کیفیت منابع آب
Groundwater Geochemistry

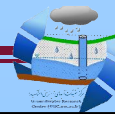
فصل پنجم

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INORGANIC CARBON — DIC

Inorganic or mineralized carbon is carbon in its most oxidized state, with a valence of 4+. Carbon dioxide is its only gaseous form, and in minerals is found coordinated with 3 oxygen atoms as carbonate. Dissolved in water, it is present in four forms, $\text{CO}_{2(\text{aq})}$, carbonic acid — H_2CO_3 , bicarbonate — HCO_3^- , and the carbonate ion — CO_3^{2-} , collectively referred to as dissolved inorganic carbon, or DIC:

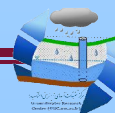
$$\text{DIC} = m_{\text{CO}_2} + m_{\text{H}_2\text{CO}_3} + m_{\text{HCO}_3^-} + m_{\text{CO}_3^{2-}}$$

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DIC can originate from a variety of reactions such as:

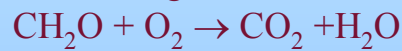
Dissolution of CO₂ from the atmosphere or from the soil:



Dissolution of calcite:



Oxidation of organics in water:



Metamorphic reactions:

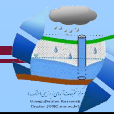


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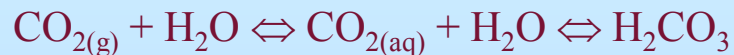
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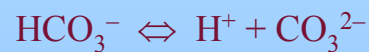
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CO₂, carbonic acid and DIC species



Carbonic acid is a weak acid, and readily dissociates to form bicarbonate, which again dissociates to form dissolved carbonate:

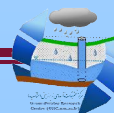


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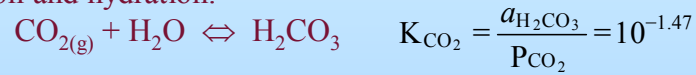
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Each step in the distribution of DIC species is controlled by a reaction constant, K, the first being the Henry's Law constant, for the dissolution of CO₂.

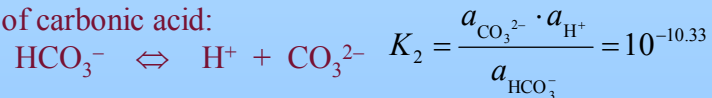
CO₂ dissolution and hydration:



1st dissociation of carbonic acid:



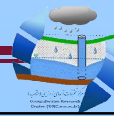
2nd dissociation of carbonic acid:



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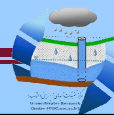
Equilibrium constants for the carbonate system.

<i>T</i> °C	<i>log K</i> _{CO2}	<i>log K</i> ₁	<i>log K</i> ₂	<i>log K</i> _{CaCO3}
0	-1.11	-6.58	-10.62	-8.38
5	-1.19	-6.52	-10.55	-8.39
10	-1.27	-6.46	-10.49	-8.41
15	-1.34	-6.42	-10.43	-8.43
20	-1.41	-6.38	-10.38	-8.45
25	-1.47	-6.35	-10.33	-8.48
50	-1.72	-6.28	-10.16	-8.65

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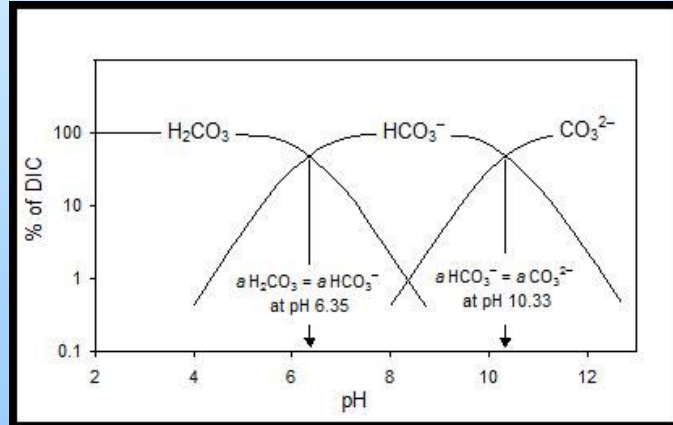
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pH and distribution of carbonate species

Relative distribution of DIC species in pure water as a function of pH, at 25°C.

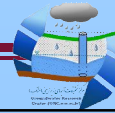


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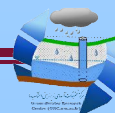
EXAMPLE: Determine the relative concentrations of the principal DIC species in water for pH 7.0 and for pH 10.0.

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Alkalinity

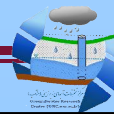
The dissociation of carbonic acid provides protons, H^+ , that are essential for the weathering of carbonate and silicate minerals. These mineral dissolution reactions act to increase both the pH and the concentration of bicarbonate $[HCO_3^-]$ and $[CO_3^{2-}]$ in solution. From the above reactions, these species are capable of consuming H^+ ions, and so are important in buffering acidity in solution. This is the attribute of alkalinity—the capacity of water to buffer pH during the addition of acid. In most waters, the species that contributes the most to alkalinity is bicarbonate, by conversion to carbonic acid H_2CO_3 :



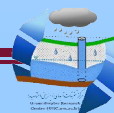
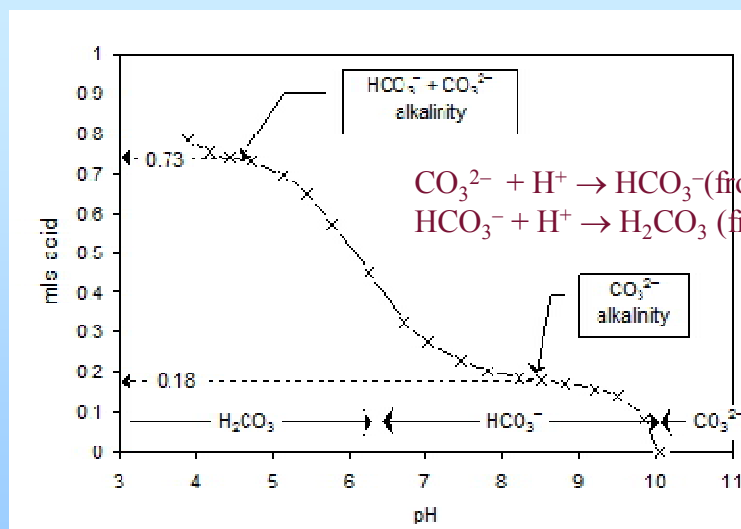
$$\text{Total alkalinity} = m_{HCO_3^-} + 2m_{CO_3^{2-}} + m_{H_3SiO_4^-} + m_{HS^-} + m_{H_2BO_3^-} + m_{OH^-} - m_{H^+}$$

$$\text{Carbonate alkalinity} = m_{HCO_3^-} + 2m_{CO_3^{2-}}$$

$$\text{Carbonate alkalinity} = \text{Total alkalinity} - m_{H_3SiO_4^-} - m_{HS^-} - m_{H_2BO_3^-} - m_{OH^-}$$



Alkalinity titrations



$$\text{Alkalinity (moles/L)} = \frac{(\text{volume of acid}) \times (\text{normality of acid})}{(\text{volume of sample})}$$

$$M_{\text{CO}_3^{2-}} = \frac{0.18 \text{ mls} \times 0.16\text{N}}{100 \text{ mls}} = 2.9 \cdot 10^{-4} \text{ mol/L} = 0.29 \text{ mmol/L}$$

$$M_{\text{HCO}_3^-} + 2M_{\text{CO}_3^{2-}} = \text{total alkalinity} = \frac{0.73 \text{ mls} \times 0.16\text{N}}{100 \text{ mls}} = 0.00117 \text{ mol/L} = 1.17 \text{ mmol/L}$$

$$M_{\text{HCO}_3^-} = \text{total alkalinity} - 2M_{\text{CO}_3^{2-}} = 1.17 - 2 \times 0.29 = 0.59 \text{ mmol/L}$$

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